

Lowland catchment runoff response to climate change under CMIP6 in the Baltic region

Darius Jakimavičius*, Diana Šarauskienė, Jūratė Kriauciūnienė, Aldona Jurgelėnaitė

Abstract

Record or near-record high or low river flows are more often observed in different regions of the world. A thriving society must understand the magnitude of these changes in the future, mitigate their negative impacts, and be prepared to live in a different world. That is why qualified, constantly updated scientific projections of future changes are essential. Neither Lithuania nor the other Baltic countries have yet assessed runoff changes according to the latest climate change projection tools outlined in the IPCC 6th AR on climate change. In this study, the HBV model was used to project potential changes in river runoff. The ranking procedure was developed and used to select the best-fit GCMs that most accurately reproduced the climate conditions of Lithuania. Due to the anticipated changes in climatic factors affecting the studied rivers, the average annual discharge is projected to decrease by 12 to 42%, depending on the hydrological region (i.e., the conditions of river runoff formation) and the selected future period. High flows (Q5) are likely to decline very similarly to the annual ones, while low flows (Q95) are expected to decrease by approximately two-thirds compared to the reference period. An uncertainty analysis of the projections revealed that GCMs contributed up to two-thirds of the total uncertainty in the final results.

Keywords

River runoff; Climate change; CMIP6; HBV

Laboratory of Hydrology, Lithuanian Energy Institute, Breslaujos St. 3, LT-44403 Kaunas, Lithuania

*Correspondence: darius.jakimavicius@lei.lt (D. Jakimavičius)

Received: 10 April 2025; revised: 8 November 2025; accepted: 12 November 2025

1. Introduction

The sustainable development of human society and the prosperity of all living organisms are highly dependent on the availability of water resources. In achieving the Sustainable Development Goals, water is positioned at front-and-center in the water-energy-food nexus systems (Susnik et al., 2023). An overwhelming amount of scientific evidence indicates the detrimental impact of ongoing changes on irreversible processes in planet ecosystems (Bongaarts, 2019; Lennox et al., 2019; Dialogue Earth, 2022; IPCC, 2023a). Freshwater ecosystems are particularly vulnerable to human-induced climate change because (i) their species have limited dispersal potential as the environment changes, (ii) water temperature and availability depend on climate, and (iii) many of these systems are already exposed to multiple anthropogenic stressors (Woodward et al., 2010). Rivers are essential providers of ecosystem services; therefore, understanding how climate change affects river hydrological processes is crucial (Yeakley et al., 2016; Etukudoh et al., 2024).

According to the most comprehensive climate change analysis published in the IPCC Sixth Assessment Report (AR6) (Calvin et al., 2023), there are no clear trends in changing streamflow at the global level. However, regional trends do emerge, with a generally increasing trend in the northern high-latitude regions and mixed trends in the rest of the world. Researchers are constantly looking for a regularity or pattern that may help them understand the processes. The runoff formation process is very complex, and even well-established hypotheses such as the DDWW (dry regions get drier and wet regions wetter) paradigm (Held and Soden, 2006), which explains many tendencies, are challenged by both observational data and modeling studies (Yang et al., 2019; Xiong et al., 2022).

Strong deviations in river flow from long-term historical patterns manifest as floods and droughts, posing challenges to households, public health, agriculture, energy, transportation sectors, and many other vital sectors of human life. The most recent report on the European State of the Climate, released jointly by the Copernicus Climate Change Service and the World Meteorological Organization (C3S, 2024), states that since the 1980s, Europe has warmed at twice the global average rate, making it the

44 fastest-warming continent on Earth. Key findings include
45 that nearly a quarter of the river network experienced
46 "exceptionally high" flows in December. Record or
47 near-record discharges were observed in major river catch-
48 ments, including the Loire, Rhine, and Danube, largely
49 due to a series of storms from October to December. In
50 contrast, drought conditions were reported in catchments
51 such as the Ebro, which had near-record low discharges in
52 May, and the Po, which experienced below-average flows
53 throughout the year, with near-record lows from Febru-
54 ary to April (C3S, 2024). Facing such dramatic changes,
55 it is crucial to understand how altered river flows and
56 their extremes may evolve in the future as the potential
57 costs of inaction may be enormous. According to the Euro-
58 pean Environmental Agency (EEA, 2023), between 1980
59 and 2022, weather- and climate-related extreme events
60 caused economic losses of assets estimated at EUR 650
61 billion in the EU Member States, of which EUR 59.4 bil-
62 lion occurred in 2021 and EUR 52.3 billion in 2022. Es-
63 timates show that each additional 0.5°C of warming in
64 China alone is projected to increase flood-related losses
65 by \$67 billion, on average (Jiang et al., 2020). Therefore,
66 growing concerns worldwide compel us to take action to
67 increase resilience and adaptability to future changes. To
68 ensure a sustainable approach to water systems
69 management, the impacts of projected climate change must
70 be understood and incorporated into regional water man-
71 agement strategies (Döll et al., 2015). That is why, living
72 in such a rapidly changing environment, decision-makers
73 need reliable and up-to-date projections of changes
74 in the hydrological regime, along with assessments
75 of the associated uncertainties (Lane and Kay,
76 2021).

77 With increasing data and research experience, sci-
78 entists are rushing to improve and update climate change pro-
79 jections and periodically undertake large-scale model com-
80 parisons with the latest and most sophisticated models to
81 better understand the climate system's response to a range
82 of potential emission or concentration scenarios (Mein-
83 shausen et al., 2020). The IPCC Sixth Assessment Report
84 (IPCC, 2023b) that gives the most complete information
85 available on the subject to date has been called the stark-
86 est warning yet about unprecedented global changes (The
87 Guardian, 2021). Along with the latest IPCC report, new
88 state-of-the-art global climate models, known as CMIP6
89 models ([Coupled Model Intercomparison Projects](#)), were
90 released. In addition, scenarios from CMIP5, known as Rep-
91 resentative Concentration Pathways, were replaced with
92 a new range of scenarios based on Shared Socioeconomic
93 Pathways (SSPs) (IPCC, 2023a).

94 Research on river runoff projections is evolving in par-
95 allel with growing knowledge of global climate change. Us-
96 ing hydrological models enhances climate change impact
97 assessments by capturing the spatial and seasonal vari-
98 ability in hydrological responses (Piniewski et al., 2018).

99 Each study typically relies on the most up-to-date genera-
100 tion of climate scenarios and hydrological models, which
101 are selected according to individual criteria (Clark et al.,
102 2017). In Lithuania, projections of river runoff were based
103 on the Special Report on Emissions Scenarios (SRES) (Kri-
104 aučiūnienė et al., 2008; Kriaučiūnienė et al., 2013) and the
105 Representative Concentration Pathways scenarios
106 (Stonevičius et al., 2017; Šarauskienė et al., 2018; Kri-
107 aučiūnienė et al., 2019; Jakimavičius et al., 2020; Akst-
108 inas et al., 2020). To date, the studies mentioned above
109 for runoff predictions have used regional climate mod-
110 els (RCMs). However, according to the IPCC AR6, only
111 global climate models (GCMs) are currently available. The
112 present study was designed to determine the effect of
113 a changing climate on Lithuanian lowland river runoff ac-
114 cording to CMIP6-based GCMs. Many studies have shown
115 that GCMs are the most versatile and effective tools for cre-
116 ating possible future climate scenarios (Bian et al., 2021).
117 Each release of a new suite of GCMs (Eyring et al., 2016),
118 updated with the latest findings, provides an opportunity
119 to reassess the impact of a changing climate on the envi-
120 ronment and society. Because the performance of GCMs is
121 site-specific, researchers in different countries employ dif-
122 ferent procedures to select those that work best for the
123 country or region they are studying. Accomplished stud-
124 ies in different countries reveal different best-performing
125 GCMs with respect to temperature or precipitation indices
126 (Raju and Kumar, 2020; Iqbal et al., 2021; Kurniadi et al.,
127 2023; Nguyen-Duy et al., 2023; Rivera, 2024; Anil et al.,
128 2024; Bhanage et al., 2024; Tariq et al., 2024).
129 However, a major limitation of global climate models is
130 their coarse spatial resolution, typically exceeding
131 $1^\circ \times 1^\circ$, which is insufficient to capture local climatic fac-
132 tors that govern river runoff formation. Therefore, the
133 research team (Gebrechorkos et al., 2023) reduced the grid
134 size to 0.25° using statistical downscaling. The
135 global climate models modified in this way could be
136 used to simulate runoff projections in lowland river
137 catchments.

138 Neither Lithuania nor the other Baltic countries have
139 evaluated runoff changes according to the newest climate
140 change research tools presented in the IPCC AR6. There-
141 fore, this study examines the impact of climate change on
142 lowland river runoff for the first time by using SSPs and
143 new GCMs. From a set of 18 models, three GCMs that best
144 correspond to the natural conditions of Lithuania were
145 selected based on a proposed ranking procedure. This
146 work will generate fresh insight into potential changes in
147 average and extreme river discharge values in the near and
148 far future for lowland rivers. Uncertainties in river runoff
149 projections arising from the selected climate scenarios and
150 global climate models will be assessed. In the absence of
151 regional climate models, the developed methodology for
152 applying global climate models could be effectively used
153 for other lowland catchments.

2. Materials and methods

2.1 Study area and data

The objects of this study are the Nemunas River and its major tributaries: Merkys, Neris, Nevezis, Dubysa, Šešupė, Jūra, and Minija (Figure 1). The main characteristics of the rivers included in the hydrological modeling are presented in Table 1. As the water gauging stations of these rivers are located at elevations of up to 78 meters above sea level, the rivers are classified as lowland rivers. The area of the Nemunas catchment at its mouth is 97,928 km², and an average discharge into the Curonian Lagoon is 605 m³ s⁻¹. The areas of the Nemunas sub-catchments range from 1,220 to 24,500 km², with average discharges varying between 14 and 160 m³ s⁻¹. Table 1 also presents the feeding sources of the studied rivers and the seasonal distribution of runoff (expressed as a percentage of the annual runoff). In the studied region, river runoff is formed

by groundwater, snow, and rainfall (Akstinas et al., 2022). Groundwater supply is represented by G, snow by S, and rainfall by R. The dominant feeding source is indicated by a capital letter, while the following feeding sources are marked with lowercase letters. For example, if groundwater is the dominant source of river runoff, while snow and rainfall contribute a smaller portion, it is marked as G-sr. The distribution of runoff throughout the year was studied over three periods previously proposed by Gailiušis et al. (2001).

For the development of hydrological models, daily precipitation (P, mm) and air temperature (T, °C) data from 14 meteorological stations (MS) (1. Dotnuva, 2. Kaunas, 3. Klaipėda, 4. Laukuva, 5. Lazdijai, 6. Panevėžys, 7. Raseiniai, 8. Šilutė, 9. Šiauliai, 10. Telšiai, 11. Ukmergė, 12. Utēna, 13. Varėna, and 14. Vilnius) as well as daily discharges (Q, m³ s⁻¹) from 11 water gauging stations (WGS) (1. Nemunas-Druskininkai, 2. Merkys-Puvočiai, 3.

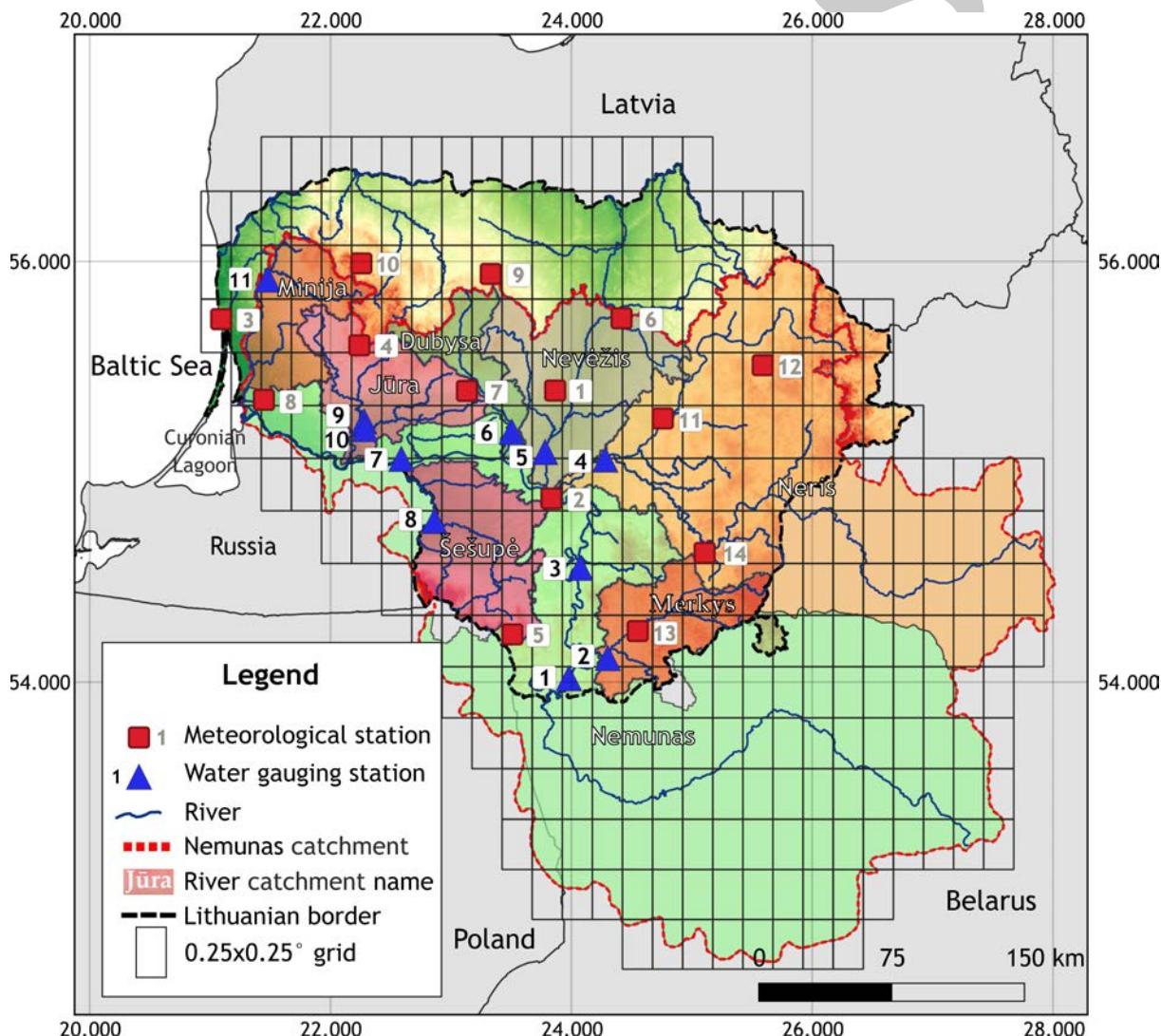


Figure 1. The Nemunas River catchment and subcatchments, meteorological and water gauging stations.

Table 1. Main characteristics of the selected rivers (according to the data in the reference period).

River-WGS	Catchment area, km ²	Altitude of WGS, m.a.s.l.	Feeding source	Q, m ³ s ⁻¹			Seasonal distribution of runoff, %		
				Annual	High flow (Q5)	Low flow (Q95)	Spring (March–April)	Summer (May–August)	Autumn (September–February)
Nemunas-Druskininkai	37400	77.49	G-sr	198	385	101	28.5	28.2	43.3
Merkys-Puvočiai	4300	78.05	G-rs	32.1	52.7	19.5	23.0	30.2	46.8
Nemunas-Nemajūnai	42900	50.65	G-sr	240	461	127	27.8	28.8	43.4
Neris-Jonava	24500	34.12	G-sr	160	318	79	28.2	27.2	44.6
Nevėžis-Babtai	5780	17.54	S-rg	31.5	115	3.54	37.1	15.2	47.7
Dubysa-Padubysis	1840	28.97	R-sg	14.2	43.6	2.99	29.0	18.0	53.0
Nemunas-Smalininkai	81200	7.33	G-sr	478	970	246	28.4	27.1	44.5
Šešupė-K.Naumiestis	3180	26.96	R-sg	34.0	102	6.12	34.0	19.7	46.3
Jūra at the mouth	3994	6.50	R-sg	42.6	162	6.13	24.1	13.6	62.3
Minija-Kartena	1220	18.01	R-sg	16.9	61.5	2.11	23.5	13.2	63.3
Nemunas at the mouth	97928	0.11	G-sr	605	1318	276	28.4	27.1	44.5

189 Nemunas-Nemajūnai, 4. Neris-Jonava, 5. Nevėžis-Babtai, 6.
 190 Dubysa-Padubysis, 7. Nemunas-Smalininkai, 8. Šešupė-K.
 191 Naumiestis, 9. Jūra-Tauragė, 10. Šešuvė-Skirgailai, and
 192 11. Minija-Kartena) for the period 1995–2014 were used
 193 (Figure 1). This 20-year period was selected in accordance
 194 with the IPCC AR6 recommendations (Calvin et al., 2023).
 195 The data mentioned above were obtained from the hydro-
 196 logical and meteorological yearbooks of the Lithuanian
 197 Hydrometeorological Service under the Ministry of Envi-
 198 ronment. Discharge projections were based on data (P
 199 and T) from global climate models that had already been
 200 statistically downscaled applying the bias correction con-
 201 structed analogues with quantile mapping reordering (BC-
 202 CAQ) method (Gebrechorkos et al., 2023). These data
 203 are freely available for scientific purposes in the CEDA
 204 (Centre for Environmental Data Analysis) database. To
 205 identify climate models that adequately represent the cli-
 206 matic conditions of Lithuania, P and T data from all 18
 207 models in the CEDA database for 1995–2014 were ana-
 208 lyzed. Using a ranking method (section 2.2.1), three cli-
 209 mate models were chosen. Their P and T data, under the
 210 SSP245 and SSP585 scenarios, were then applied to project
 211 river discharge for the near (2031–2050) and far future
 212 (2081–2100).

2.2 Methodology

214 The assessment of changes in the Lithuanian rivers' runoff
 215 according to SSP scenarios and global climate models was
 216 carried out in four stages. In the first stage, three out of 18
 217 global climate models (already statistically downscaled)
 218 that most accurately represent the climatic conditions of
 219 Lithuania were selected. In the second stage, hydrological
 220 models of the rivers were developed, calibrated, and vali-
 221 dated. In the third stage, river discharge was simulated for
 222 the near and far future using the selected climate models
 223 and the created hydrological models, under the two most

224 commonly applied SSP scenarios (SSP245 and SSP585). In
 225 the fourth and final stage, the contributions of global cli-
 226 mate models (GCM) and SSP scenarios to the uncertainties
 227 in runoff projections were quantified.

2.2.1 Climate model selection

228 Hydrological modeling based on the output data of GCMs
 229 is often used to assess future changes in river runoff. The
 230 most recent SSP scenarios and GCMs, proposed by the Sixth
 231 Assessment Report (IPCC, the Sixth Assessment Report
 232 (AR6)), are currently being applied. A large amount of data
 233 (P and T) from these climate models is available in open-
 234 source databases. However, in river hydrological modeling,
 235 it is important to choose models that accurately represent
 236 the climatic conditions of the study area. In practice, two
 237 approaches are commonly applied for this purpose: 1)
 238 based on the output data of all available GCMs, the median,
 239 lower and upper limit of the applied ensemble are derived,
 240 and 2) all available GCMs are used to simulate the past
 241 climate conditions and the best-performing GCM is then
 242 selected. The first approach provides a broad spectrum of
 243 future climate parameters, which will not always precisely
 244 capture the local climatic patterns. Meanwhile, the second
 245 approach is based on the assumption that climate models
 246 capable of reproducing the past climate with satisfactory
 247 accuracy are likely to provide more reliable projections of
 248 future conditions. Therefore, it was decided to apply the
 249 second approach, i.e., to select three climate models and
 250 use their average output data to project the discharge of
 251 rivers in the Nemunas catchment for the near (2031–2050)
 252 and far (2081–2100) future.

253 Five parameters were used for model selection: daily
 254 Q-Q plot, monthly standard deviation, and the minimal,
 255 average, and maximum values of precipitation and tem-
 256 perature. All five parameters were assigned equal weights
 257 because, in the absence of prior information favoring any

Table 2. Calibration and validation results of hydrological models.

Subcatchments	Calibration			Validation		
	r	NSE	RE, %	r	NSE	RE, %
Nemunas at Druskininkai	0.85	0.69	-1.48	0.76	0.50	1.40
Merkys	0.81	0.54	0.23	0.82	0.59	-0.20
Nemunas at Nemajūnai	0.85	0.69	-2.28	0.76	0.51	2.18
Neris	0.84	0.61	4.16	0.85	0.59	-3.76
Nevėžis	0.84	0.70	-0.70	0.76	0.50	1.31
Dubysa	0.85	0.73	-0.41	0.78	0.60	0.43
Nemunas at Smalininkai	0.86	0.72	-0.69	0.79	0.52	0.71
Šešupė	0.89	0.79	2.61	0.76	0.51	-2.14
Jūra	0.87	0.74	4.14	0.86	0.73	-3.85
Minija	0.85	0.72	-1.88	0.84	0.70	1.83
Nemunas at the mouth	0.90	0.81	0.12	0.83	0.60	0.12

Table 3. Summary of global climate model ranking results.

Models	Precipitation, P					Air temperature, T					SUM	RANK
	Q-Q plot	Average	STDEV	MIN	MAX	Q-Q plot	Average	STDEV	MIN	MAX		
ACCESS-CM2	143	146	59	131	117	107	106	125	213	156	1303	10
BCC-CSM2-MR	99	98	141	120	124	113	113	129	73	173	1183	4
CESM2	98	99	217	118	210	184	181	129	76	161	1473	14
CMCC-CM2-SR5	101	98	112	124	108	147	147	107	83	89	1116	2
CMCC-ESM2	107	109	94	152	123	83	81	155	232	158	1294	8
GFDL-ESM4	147	145	142	103	128	122	122	156	126	160	1351	11
HadGEM3-GC31-LL	168	168	186	128	177	215	214	200	76	100	1632	18
IITM-ESM	205	212	92	149	127	128	126	111	172	97	1419	12
INM-CM4-8	161	160	142	83	131	177	178	114	121	154	1421	13
INM-CM5-0	139	137	202	105	151	96	98	124	91	152	1295	9
IPSL-CM6A-LR	96	94	165	190	115	139	139	121	59	135	1253	6
KACE-1-0-G	112	111	115	99	126	84	83	115	113	209	1167	3
MIROC-ES2L	97	97	67	161	107	113	113	112	100	61	1028	1
MIROC6	123	121	94	169	107	163	164	138	97	91	1267	7
MPI-ESM1-2-LR	163	160	123	135	140	138	138	120	250	117	1484	15
MRI-ESM2-0	170	167	90	163	117	192	195	187	102	105	1488	16
NorESM2-MM	167	170	229	166	155	96	97	107	192	192	1571	17
UKESM1-0-LL	98	102	124	98	131	97	99	144	218	84	1195	5

specific parameter, equal weighting was considered a neutral and unbiased approach. Consequently, each parameter contributed equally to the overall ranking of the GCMs, with no single parameter regarded as more influential than the others. Following the recommendations of AR6, model performance was assessed against observations for the period 1995–2014. In the first step, the five parameters were calculated from observational data at 14 MSs. In the second step, the same parameters for the same 14 MSs were calculated based on the outputs of 18 GCMs. In the third step, the values obtained from the observational data were compared with those derived from the outputs of GCMs. The climate model, according to the data of which a specific parameter value calculated for a specific MS was the closest to the observational one, was assigned a rank of 1, the second most similar a rank of 2, the third a rank of

3, etc. In the fourth step, the ranks of the five precipitation indicators and the five air temperature indicators were summed up. The model with the lowest total rank over all 14 MSs was considered the most suitable for the studied area, followed by the second lowest, and so on.

2.2.2 Discharge projection of the Nemunas River catchment using the HBV hydrological model

The HBV (*Hydrologiska byråns vattenbalansavdelning*) model, developed at the Swedish Meteorological and Hydrological Institute (Bergstrom, 1992), was used to project the runoff of the Nemunas River catchment according to global climate models and Shared Socioeconomic Pathway scenarios. This hydrological model is widely used to address the impact of climate change on river hydrology (Pervin et al., 2021). Even though this software was originally developed in the early 1970s, it has undergone

continuous improvements. This model requires relatively limited input data, including precipitation, air temperature, and geographical information of the river catchment for which runoff is modeled (catchment area, height above sea level, forest cover, lake cover, MS-defined catchment area). Due to its relative simplicity, various versions of the HBV model have been applied in more than 30 countries across diverse climatic conditions, e.g., Sweden, Zimbabwe, India, Colombia (Bergstrom, 1992). The HBV has also been successfully applied in our previous studies (Jakimavičius et al., 2018; Akstinas et al., 2020).

The model calculations were performed in three steps. In the first step, the amount of precipitation that reaches the ground was estimated. In the second, slope runoff was simulated; and in the third, river discharge and its transformation within the watercourse were evaluated.

The HBV model is based on the water balance equation (IHMS, 2005):

$$P - E - Q = \frac{d}{dt} [SP + SM + UZ + LZ + V] \quad (1)$$

where P is precipitation, E is evaporation, Q is discharge, SM is soil moisture, SP is snowpack, UZ is upper ground water zone, LZ is lower groundwater zone, and V is lake or dam volume.

For the development of the Nemunas River hydrological model, daily discharge data from 11 WGSs, as well as air temperature and precipitation data from 14 MSs, were used (Figure 1). The same information about the modeled catchment area, the presence of lakes and forests as well as mean elevation above sea level was used for both the reference period and the projections. The hydrological model consisted of 11 subcatchments: the Nemunas at Druskininkai, Nemajūnai, Smalininkai, and its mouth, together with its main tributaries in sequence – the Merkys, Neris, Nevezis, Dubysa, Šešupė, Jūra, and Minija (Figure 1).

Following the recommendations of IPCC AR6 (Calvin et al., 2023), the period from 1995 to 2004 was selected for model calibration, whereas the period 2005–2014 was used for validation. The calibration procedure involved adjusting 16 model parameters and comparing calculated discharge values with the observed ones. Four groups of calibration parameters were used in the calibration process (IHMS, 2005): the model parameters that (i) control general runoff volume over the total calibration period, (ii) describe snow accumulation and melting intensity, (iii) characterize the moisture accumulated in soil, and (iv) define the extremes (river floods and droughts) in discharge hydrograms. During the spring flood, the most important calibration parameters for runoff modeling are related to snowmelt and soil moisture storage, while during the low water period - the parameter that determines river underground feeding (Kriauciūnienė et al., 2013).

Ideally, the correlation coefficient (r) should approach 1; however, values above 0.7 are considered acceptable for proper calibration (IHMS, 2005). Similarly, the hydrological model can be regarded as calibrated when the Nash–Sutcliffe efficiency (NSE) exceeds 0.5 (Ritter and Muñoz-Carpena, 2013). Calibration and validation results for each subcatchment are presented in Table 2. Based on the obtained values of r , NSE, and RE (difference between observed and calculated discharge), it was decided that the hydrological model of the Nemunas River is ready to perform discharge projections for the near and far future using climate models data.

2.2.3 Estimation of uncertainty sources in projected runoff of the Nemunas River catchment

This study considered uncertainties in runoff projections arising from the selection of global climate models and SSP scenarios. All possible combinations (24 combinations for each of the eight rivers, i.e. (3 GCMs \times 2 scenarios + 2 scenarios \times 3 GCMs) \times 2 periods (near and far future)) of uncertainty sources were analyzed to identify the two main sources of uncertainty. The assessment was conducted in four steps: 1) river discharge was calculated for each model and SSP scenario for the near and far future; 2) the differences between the lowest and highest water discharges under GCMs or SSP scenarios were estimated separately for each period; 3) the average of discharge differences was calculated for the scenarios and GCMs separately in the near (2031–2050) and far (2081–2100) future periods; 4) the relative contribution (%) of each model and scenario to the overall uncertainty was quantified based on these discharge differences.

For comparison, three models with the lowest ranking scores (Figure 2b-d) and three with the highest scores are presented (Figure 2e-g). A visual assessment revealed that the distributions of the lowest-ranking models differed only slightly from the distribution derived from observational data. Therefore, we assumed that if these models were able to reproduce past climate conditions with sufficient accuracy, then their future predictions should be suitable for the assessment of the climate conditions in the studied region.

3. Results

3.1 Climate model selections

Based on the methodology presented in section 2.2.1, daily Q-Q plots, monthly standard deviations (STDEV), minimal, average, and maximum values (of P and T) were calculated for 14 MSs using both observed data and output data from 18 GCMs. The results were arranged and summarized over all MSs (Table 3). The ranking results showed that the applied models received different scores depending on the evaluation criteria. If considering only the accuracy of P projections, the IPSL-CM6A-LR model exhibited the smallest deviations from the actual data according

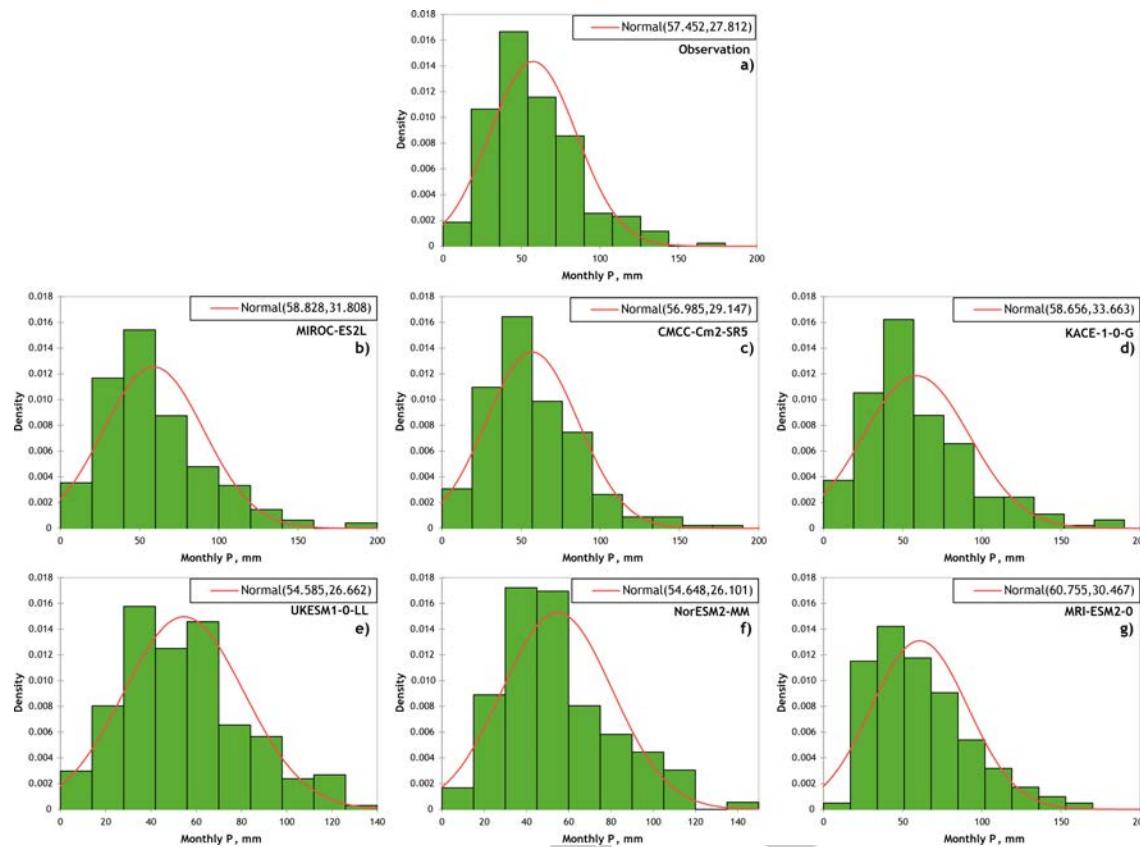


Figure 2. Comparison of observed (a) precipitation distributions with global climate models that scored the lowest (b-d) and highest (e-g) ranks.

to the Q-Q area and the average monthly P. The ACCESS-CM2 distinguished itself in terms of STDEV. Meanwhile, the INM-CM4-8, MIROC-ES2L, and MIROC6 models stood out when evaluating the monthly min and max P, respectively. The evaluation of T projections yielded somewhat different results. The CMCC-ESM2 model demonstrated the best performance in terms of the Q-Q area and average monthly T. Considering the average monthly T variability (STDEV), the CMCC-CM2-SR5 and NorESM2-MM got equal ranks. The minimum T was most accurately projected by the IPSL-CM6A-LR, and the maximum by the MIROC-ES2L model.

The river discharge was projected by applying P and T data according to various scenarios. Therefore, the selected GCMs (or their ensemble) must provide the most accurate possible estimates of both indicators. After summarizing all ranking criteria, we found that in the historical period, STDEV of P and T from actual observations were the smallest in the case of three models: MIROC-ES2L, CMCC-CM2-SR5, and KACE-1-0-G. Additionally, a visual comparison of the distributions of the selected models' outputs was performed. The distribution of average monthly P data was compiled based on the observation data from 14 MSs in the reference period (Figure 2a).

3.2 Changes in the conditions of runoff formation in the Nemunas River catchment according to global climate models and SSP scenarios

The runoff of Lithuanian rivers is shaped by physical-geographical and climatic conditions. Based on regional differences in these conditions, three hydrological regions are distinguished: western (W-LT), central (C-LT), and south-eastern (SE-LT) (Akstinas et al., 2022) (Figure 3). In W-LT, the greatest amount of precipitation falls. Combined with steep river slopes and favorable conditions for rapid water flow, this results in rivers being predominantly rain-fed, with rainfall accounting for 62% of their total runoff. In C-LT, river slopes are small, and impermeable soils are widespread, which creates more favorable conditions for evaporation. Summer precipitation is low, and the underground supply is scarce (17%), so rivers become even more depleted. In SE-LT, the relief is gradually rising, which increases river slopes. Runoff in this region is determined by a higher amount of precipitation compared to C-LT and abundant underground feeding (55%). Due to the reasons above, rivers here carry more water than in C-LT but are less watery than in W-LT.

Before analyzing future changes in river discharge, it should be helpful to find out how runoff formation conditions would change under the applied GCMs and SSPs

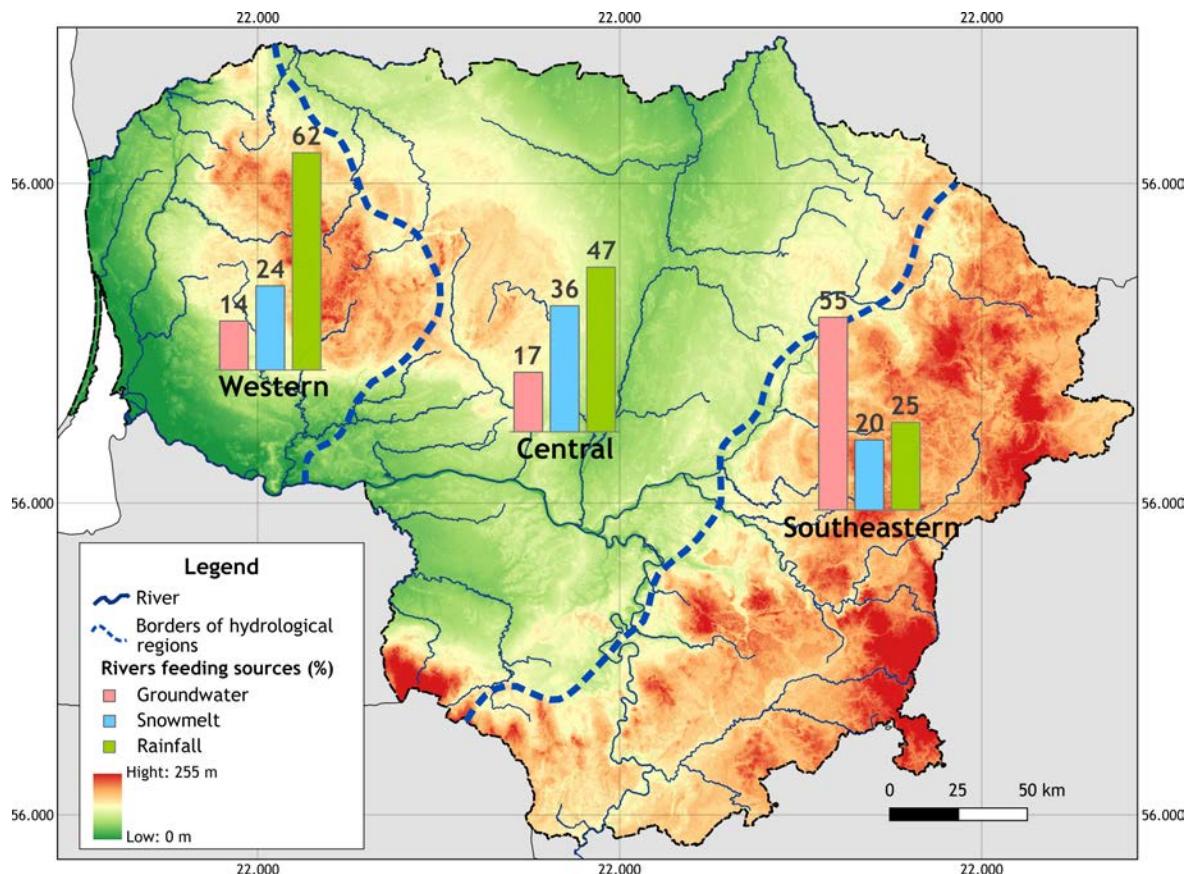


Figure 3. Hydrological regions of Lithuania (based on Akstinas et al., 2022).

445 across the different hydrological regions.

446 It was determined that P and T would change considerably in the future. As shown in Figure 4, across all hydrological regions, projected changes in T are going to be 447 very similar. Based on the average of three models, the 448 mean annual T during the reference period (1995–2014) 449 was 7.4°C in W-LT, 7.2°C in C-LT, and 6.9°C in SE-LT. In the 450 near future (2031–2050), no significant differences were 451 identified between the applied scenarios. Under the most 452 likely SSP245 scenario, T would rise by 2.2–2.3°C in the 453 studied hydrological regions, and under the pessimistic 454 SSP585 scenario by 2.5–2.8°C compared to the reference 455 period. Considerably larger differences are possible in the 456 far future (2081–2100): under SSP245 scenario, T would 457 increase by 3.5–3.8°C, whereas, under the SSP585, from 458 5.8°C to 6.4°C depending on the hydrological region. 459

460 Analysis of seasonal air temperature changes does not 461 indicate significant differences between the scenarios in 462 the near future. The smallest increase is projected for autumn 463 (1.8–2.4°C), a moderate rise for spring and summer 464 (2.1–3.0°C), and the largest increase for winter (2.4–3.0°C). 465 There would be no clear trends in seasonal temperature 466 rise in the far future, but there will be apparent differences 467 between scenarios. According to the SSP245 scenario, air 468 temperature (depending on the region) is likely to rise 469

470 by 3.2–3.4°C in autumn, 3.3–3.7°C in spring, 3.5–3.8°C 471 in winter, and 3.8–4.2°C in summer. Meanwhile, under 472 the SSP585 scenario, substantially greater warming is 473 expected: up to 5.9°C in spring, 6.0°C in autumn, 6.5°C in 474 winter, and 7.0°C in summer. Regional comparisons show that 475 W-LT is likely to experience the smallest changes, whereas 476 C-LT and SE-LT are projected to undergo the greatest 477 increases relative to the reference period. 478

479 Based on the data from three GCMs, in the reference 480 period, the highest amount of precipitation, 811 mm per 481 year, was determined in the river subcatchments located 482 in W-LT. In C-LT and SE-LT, P was 643 and 673 mm per 483 year, respectively. Projections made using GCMs data under 484 the SSP245 scenario revealed that in the near future, P 485 should be from 2.2% (W-LT) to 3.8% (SE-LT) higher than 486 in the reference period (Figure 4). In contrast, under the 487 SSP585 scenario, P is expected to decline slightly, by 1.1% 488 in W-LT and 1.8% in C-LT. In the far future, the most 489 pronounced positive changes are projected for W-LT (5.0% 490 under SSP245 and 3.0% under SSP585), followed by C-LT 491 (3.3% and 2.1%, respectively), while the smallest increases 492 would be in SE-LT (2.6% and 1.4%). Although the average 493 annual precipitation may change slightly, significant 494 positive and negative changes in the seasonal amount of 495 precipitation are projected. The most significant positive 496

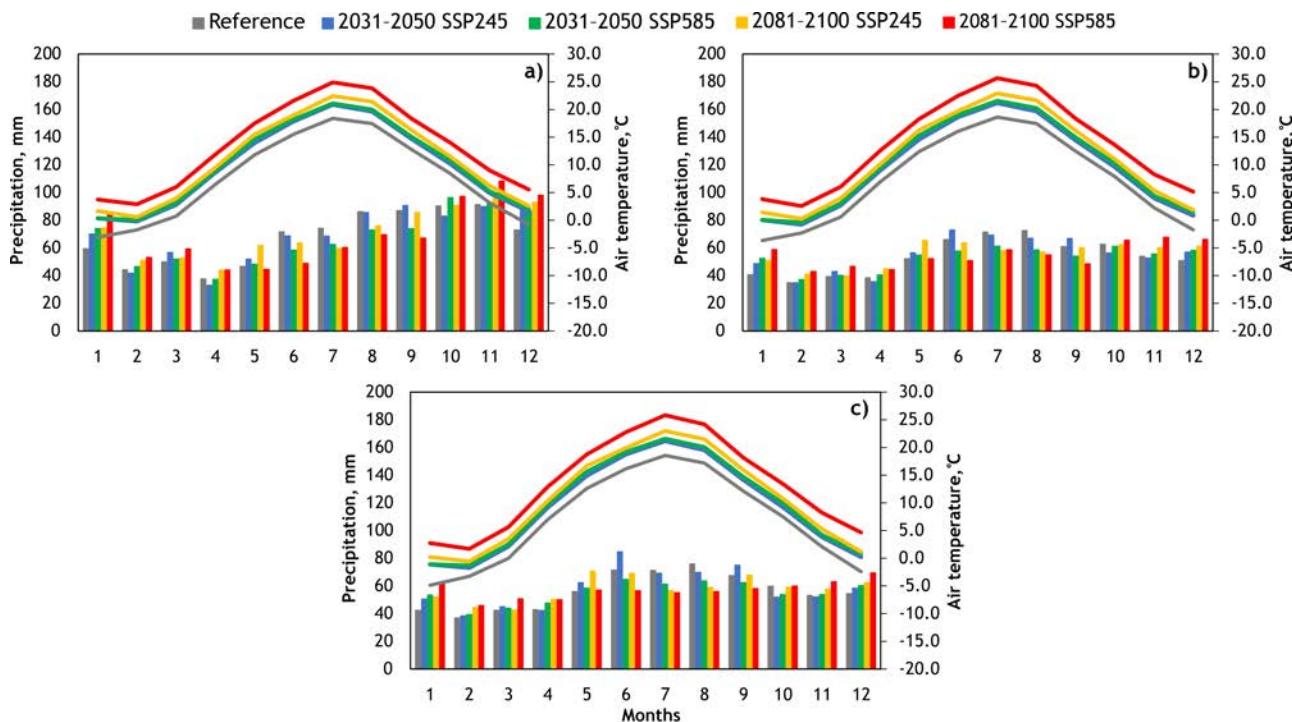


Figure 4. Projection of air temperature (°C) and precipitation (mm) in the western (a), central (b), and south-eastern (c) hydrological regions of Lithuania.

495 changes of P are expected in winter. In the near future, winter
 496 precipitation is expected to rise by 10.2–17.7%, and
 497 in the far future, by 18.9–33.7%, relative to the reference
 498 period. Spring is expected to experience smaller positive
 499 changes: precipitation is likely to increase by up to 6.3% in
 500 the near future, depending on the region and scenario, and
 501 by up to 18.0% in the far future. In contrast, significant
 502 negative changes in precipitation are projected for summer.
 503 In the near future under the SSP245 scenario, precipitation
 504 may decline by up to 3.9% in W-LT and C-LT, while
 505 SE-LT may experience an increase of up to 2.4%. However,
 506 according to the SSP585 scenario, summer precipitation
 507 would decrease from 13.2% to 16.4%, depending on the
 508 hydrological region. In the far future, the most substantial
 509 decreases are expected in summer, ranging from 14.2% to
 510 23.2%. Meanwhile, in autumn, both negative and positive
 511 changes in precipitation are expected, depending on the
 512 projection period. In the near future, the slightest negative
 513 changes would be in W-LT (up to 2.0%), more significant
 514 in C-LT (up to 3.9%), and the largest in SE-LT (up to 5.9%).
 515 However, in the far future, the amount of precipitation is
 516 expected to increase by 0.6–2.9% relative to the reference
 517 period.

518 3.3 Projections of the Nemunas River catchment dis- 519 charge in the near and far future

520 Discharge simulations for the near (2031–2050) and far
 521 (2081–2100) future were carried out using the outputs of
 522 three GCMs (MIROC-ES2L, CMCC-CM2-SR5, and KACE-1-

523 0-G) under two SSPs (SSP245 and SSP585). The results
 524 were compared with the results of discharge simulations
 525 of the same models for the reference period (1995–2014).
 526 The estimated changes in Lithuanian river discharge had
 527 different regional patterns. Therefore, the analysis was
 528 performed at two spatial scales: the entire Nemunas catch-
 529 ment and individual hydrological regions, with one repre-
 530 sentative river selected from each region (Neris River for
 531 SE-LT, Nevėžis for C-LT, and Minija for W-LT). Based on the
 532 results, projected changes in average annual, high (Q5),
 533 and low (Q95) flows were assessed for the near and far
 534 future periods.

535 The projected changes in climate parameters are likely
 536 to significantly reduce the Nemunas discharge in both stud-
 537 ied future periods (Figure 5a). In the near future, the aver-
 538 age annual discharge is projected to decrease from 15.1%
 539 to 23.5%, while in the far future, from 24.2% to 41.7%
 540 compared to the reference period (Table 4).

541 The Nemunas River catchment covers 75% of Lithuania's
 542 territory and extends across all three hydrological re-
 543 gions, resulting in diverse river feeding conditions
 544 (Figure 3). In the far future, according to the most unfavor-
 545 able scenario (SSP585), a considerable decrease in the dis-
 546 charge of the Nemunas River is expected, primarily driven
 547 by a pronounced temperature rise of 5.8–6.4°C across dif-
 548 ferent hydrological regions. Although precipitation would
 549 increase slightly (1.4–3.0%) under the SSP585 scenario,
 550 this increase would not be sufficient to significantly reduce
 551 discharge in the long term. The Neris catchment is mainly

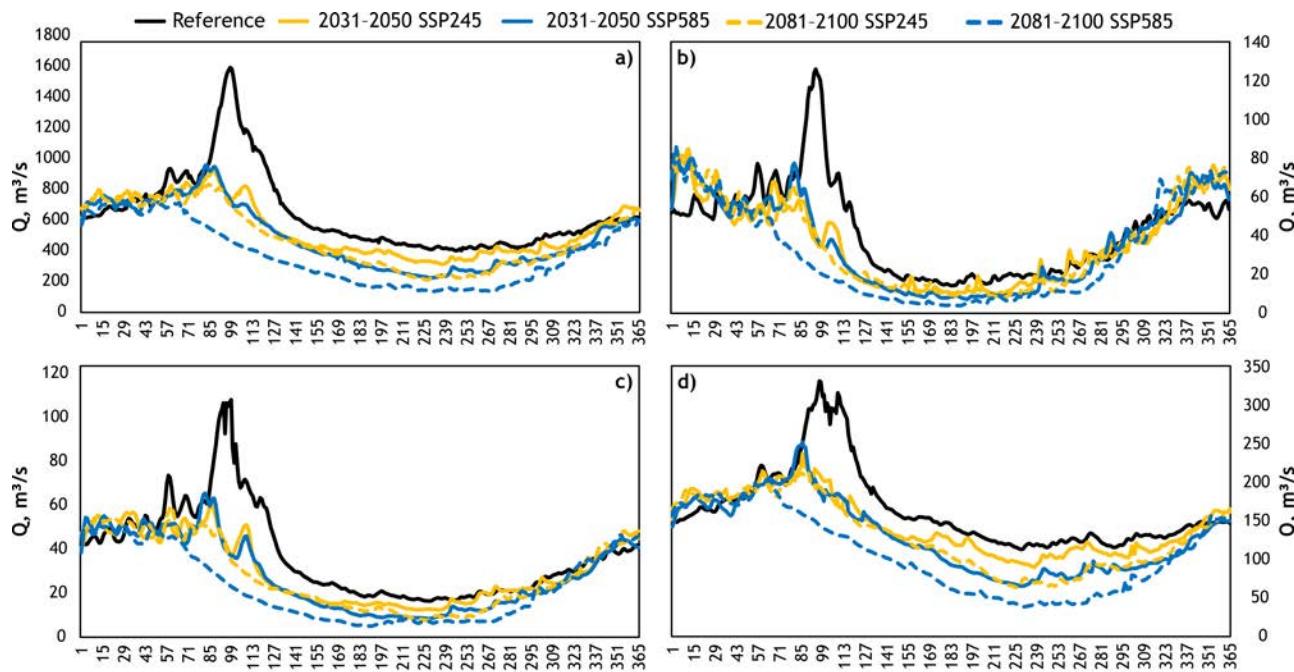


Figure 5. Nemunas (a), Minija (b), Nevėžis (c) and Neris (d) discharge projections in the near and far future compared to the reference period.

located in SE-LT, where groundwater feeding is predominant (Figure 3). As a result, the projected changes in the Neris are less notable than in the Nemunas catchment. Under both scenarios, the average annual discharge of the Neris would decline by 11.9–21.0% in the near future and by 19.9–34.7% in the far future, compared to the reference period (Figure 5d). In the Minija catchment, from W-LT, where precipitation is the primary source of river feeding (Figure 3), the discharge trend is different. Since W-LT is also projected to have more precipitation in the future, the Minija River is expected to experience the smallest reduction in discharge among the studied rivers: 14.3–15.9% in the near future and 15.3–26.7% in the far future, compared to the reference period (Figure 5b). The runoff formation of the Nevėžis River, located in C-LT, depends on both rainfall and snowmelt. However, in the far future, snowmelt floods are less likely, so its average annual discharge is projected to decrease more significantly, by 24.1–38.6% (Figure 5c).

The analysis revealed that the average annual discharges of all studied rivers are projected to change: the smallest changes are expected in the near future, while significantly larger changes are anticipated in the far future. Under the SSP245 scenario, the changes would be smaller, whereas under SSP585 they would be more pronounced. When comparing different hydrological regions, the results indicate that both in the near and far future, the smallest changes are possible in rivers from W-LT, moderate changes in rivers from SE-LT, and the largest changes in rivers from C-LT.

It was established that in all studied rivers, the high flows (Q5, typically associated with spring floods) and low flows (Q95, representing river water content during the dry season) would decrease considerably, though in different ways (Table 4). The most remarkable changes are expected in the far future when spring flood discharges (with a 5% probability) are projected to decline to a similar extent as the annual discharges. This means that floods would decline most significantly in the far future under the SSP585 scenario, as substantially higher air temperatures (especially during the winter season) are likely to prevent the formation of snowmelt-driven floods. The river discharge during the dry season (of a 95% probability) is also going to change drastically in the long term in the case of the SSP585 scenario, decreasing from 71.2% (in the Nemunas) to 81.5% (in the Minija) (Table 4). This may be due to the projected increase in summer temperatures (by up to 7.0°C) and the simultaneous reduction in precipitation compared to the reference period.

3.4 Estimation of uncertainties in the projections of the Nemunas River catchment discharge

The accuracy of river runoff projections depends on several factors, including the selected hydrological model parameters, the SSP scenario, and the global climate model (GCM). In this study, uncertainties in water flow projections were assessed only on the basis of climate models and SSP scenarios. The impact of hydrological model parameters and climate scenarios on runoff modeling results was estimated several years ago by the authors of this ar-

Table 4. Changes in discharge in the near and far future compared to the reference period.

River	Discharge	2031–2050		2081–2100	
		SSP245	SSP585	SSP245	SSP585
Nemunas	Q5	-22.3	-22.9	-28.7	-36.9
	Average	-15.1	-23.5	-24.2	-41.7
	Q95	-29.8	-47.6	-41.5	-71.2
Minija	Q5	-15.0	-12.4	-14.4	-19.9
	Average	-14.3	-15.9	-15.3	-27.6
	Q95	-44.6	-64.2	-61.3	-81.5
Nevėžis	Q5	-27.5	-27.2	-31.2	-37.2
	Average	-18.6	-23.8	-24.1	-38.6
	Q95	-36.9	-51.1	-55.1	-77.4
Neris	Q5	-16.1	-17.0	-22.1	-27.9
	Average	-11.9	-21.0	-19.9	-34.7
	Q95	-24.2	-45.7	-42.6	-73.3

SE-LT, respectively, and the rest consists of rainfall and snowmelt. Therefore, the response to climate change is more pronounced in W-LT and C-LT than in SE-LT. Even though in the near future, the influence of climate models on discharge projection results depending on hydrological regions has been clearly expressed, in the far future, these regional differences would disappear due to increasing climate extremes. Thus, in the far future, the influence of climate models should be very similar across all river sub-catchments from different hydrological regions, accounting for 63%, 64%, and 64%, respectively, with the remainder attributable to SSP scenarios.

4. Discussion

Scientific studies show that climate change strongly affects water resources, causing record high or low river flows worldwide. To adapt, society needs reliable, up-to-date scientific projections to mitigate risks and prepare for the future. To find the best-performing global climate models for projecting runoff in selected Lithuanian lowland rivers, five ranking parameters were applied: the daily Q-Q plot, monthly standard deviation, and minimal, average, and maximum values of precipitation and temperature. Eighteen GCMs from CMIP6 were ranked according to these selected parameters. Three GCMs, namely, MIROC-ES2L, CMCC-CM2-SR5, and KACE-1-0-G, received the highest scores. As the best representatives of Lithuanian climate conditions, the outputs of these models were subsequently used as inputs for hydrological simulations made for the near (2031–2050) and far (2081–2100) future periods. The general trends obtained in the recent runoff projections were quite similar to those reported in previous studies indicating a decline in spring floods and summer low flows, alongside an increase in winter discharge. However, in some cases, the scale of projected changes was greater if compared to the ones identified according to previous CMIP5 climate projections. Earlier assessments of future annual runoff revealed decreases of up to 24% (Šarauskienė et al., 2018), 31% (Jakimavičius et al., 2020), and 40% (Kriauciūnienė et al., 2019) under the most extreme scenarios in the far future. These results are consistent with the present findings showing a possible decline in this hydrological parameter from 26.7% (in the Minija) to 41.7% (in the Nemunas) under the most unfavorable scenario. Regarding dry season discharge in the far future, the present study suggests decreasing to almost 70–80% in individual catchments. In contrast, the previous findings based on CMIP5 tools indicated summer low flow reductions of only 28–43% (Šarauskienė et al., 2018). Our findings indicate that the most significant changes are expected in the central hydrological region of Lithuania, where river catchments are considered particularly sensitive to climate change. This is consistent with the results reported by other authors (Nazarenko et al., 2023). The significant negative trends in low flows observed in this

title (Kriauciūnienė et al., 2013). In both this study and the previous one, river runoff was modeled using the HBV software. For consistency, the same rivers – the Neris and the Merkys – were selected for the analysis. Therefore, the results reported by Kriauciūnienė et al. (2013) provide valuable insights into the influence of hydrological model parameters on the uncertainties in water discharge projections. That assessment showed that, for the Merkys River, the accuracy of runoff projections was determined by model parameters (7.2%), SSP scenarios (60.9%), and GCM (32%). For the Neris River, the corresponding contributions were 5.6%, 64.4%, and 30%, respectively. A previous assessment of uncertainties confirmed that, in the studied rivers, hydrological model parameters represent the smallest source of uncertainty compared with climate models or SSP scenarios. Therefore, this study assessed only the uncertainties associated with the three selected global climate models and two SSP scenarios.

The uncertainty of river projections was analyzed separately for the entire Nemunas catchment and sub-catchments representing three hydrological regions: the Minija and the Jūra in W-LT, the Šešupė, the Dubysa, and the Nevėžis in C-LT, and the Neris and the Merkys in SE-LT. The analysis revealed that in the Nemunas discharge projections for the near future, SSP scenarios and climate models had an equal impact on the final result (50% each) (Figure 6). Meanwhile, in the case of the far future, the influence of scenarios decreased to 38%, while that of climate models increased by 62%. Somewhat different regularities were established in the studied sub-catchments. In the near future, climate models accounted for 61% and 60% of the uncertainty in W-LT and C-LT, respectively, compared to 50% in SE-LT. This could be explained by differences in hydrological regimes: groundwater contributes 14%, 17%, and 55% of the discharge in W-LT, C-LT, and

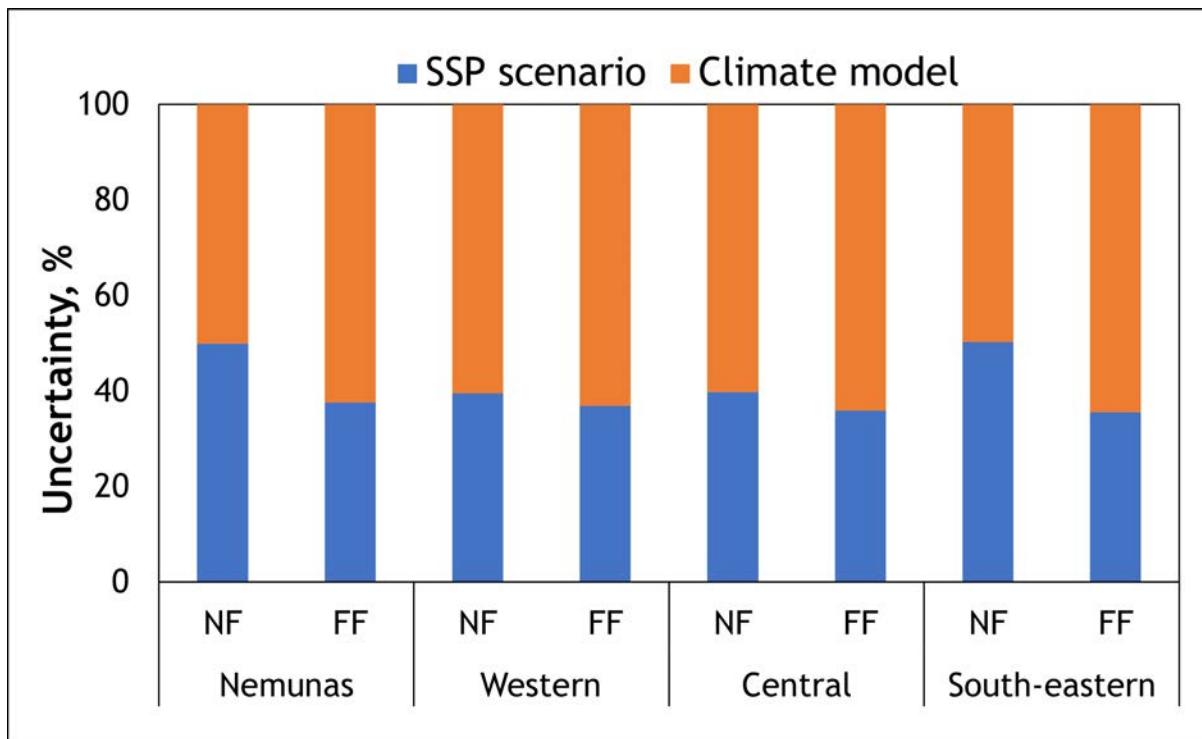


Figure 6. Uncertainties of the discharge projections in the near (NF) and far future (FF).

region in the past (Nazarenko et al., 2022), together with flow intermittency phenomena (Šarauskienė et al., 2020), indicate that this process is not new but has been ongoing for some time.

The magnitude of projected annual runoff reported in other studies exhibits considerable variation depending on the specific regions and geographical contexts examined. Such variability is observed not only in large-scale global and continental assessments (e.g., Donnelly et al., 2017; Duan et al., 2017; Yang et al., 2017; Brêda et al., 2020; Guan et al., 2021; Kis and Pongrácz, 2024) but also in more localized, national-scale investigations (e.g., Piniewski et al., 2018; Muelchi et al., 2021; Dallison et al., 2022; Murphy et al., 2023) that share methodological similarities with the present study. Interestingly, despite differences in catchment characteristics, some seasonal projections across these studies display notable similarities – for instance, increases in projected discharge during winter (Piniewski et al., 2018; Afzal et al., 2020; Muelchi et al., 2021; Slezia et al., 2021; Dallison et al., 2022; Kis and Pongrácz, 2024), decrease in summer (Afzal et al., 2020; Slezia et al., 2021; Dallison et al., 2022). One possible reason for the differences in future runoff simulation results may be the peculiarities of the CMIP6 models. Compared to the IPCC AR5, precipitation projections in the GCMs from AR6 indicate a stronger drying trend, which extends even to parts of northern Europe (Palmer et al., 2021). The higher global climate sensitivities of CMIP6 models determine higher summer temperatures in northern Eu-

rope as well. Another possible reason for the identified differences is the variation in grid resolutions among climate models. Using regional climate models (RCMs) with higher spatial resolution is recommended to obtain more accurate runoff projection results. Unfortunately, RCMs for AR6 have not yet been developed; therefore, this study used global climate models with output data already down-scaled to a $0.25^\circ \times 0.25^\circ$ grid, which may still be too coarse for catchment-scale modeling. In general, each new Phase of the Coupled Model Intercomparison Project is expected to improve model performance – just as CMIP6 GCMs are anticipated to deliver more reliable and comprehensive projections (Wei et al., 2023). Previous studies comparing the performance of CMIP6 GCMs with those from the earlier CMIP5 generation generally demonstrate an improved ability of the newest models to reproduce various temperature and precipitation patterns across different regions of the world (Chen et al., 2020; Grose et al., 2020; Gusain et al., 2020; Kim et al., 2020; Xin et al., 2020; Gebresellase et al., 2022; Martel et al., 2022; Wei et al., 2023). There is no doubt that the use of combined SSP-RCP pathways, rather than RCP emission scenarios, also influences the present results, as these pathways account for socio-economic indicators. The incorporation of Shared Socioeconomic Pathways provides a framework for accounting for potential socioeconomic developments at the global scale, thereby at least partially addressing uncertainties related to human-induced impacts on runoff conditions. Nevertheless, in this river runoff projection study, the limitations of model

758 simulations concerning human influences are unavoidable,
 759 as future modifications in catchment management, land
 760 use, hydraulic infrastructure or other anthropogenic inter-
 761 ventions cannot be reliably predicted.

762 Additionally, this study involved an uncertainty assess-
 763 ment, which is considered a very important part of the
 764 whole climate modeling process and may contribute to
 765 improving the applied modeling techniques (Tian et al.,
 766 2016; Vetter et al., 2017; Kundzewicz et al., 2018). Un-
 767 certainty ranges in the discharge projections made for
 768 the selected lowland rivers under the chosen GCMs and
 769 SSP scenarios for the near future diverged. For the rivers
 770 less dependent on precipitation, the influence of selected
 771 models and scenarios on runoff modeling results was very
 772 similar, whereas, in predominantly rain-fed and snow-fed
 773 rivers, the uncertainty attributable to GCMs accounted for
 774 60–61% of the projection results. In contrast, for the far fu-
 775 ture, the influence of the used GCMs and SSPs on the runoff
 776 projections was found to be very similar for all rivers, with
 777 more than 60% of the uncertainty from GCMs and the re-
 778 mainder from SSPs. Even though the influence of river
 779 feeding characteristics on discharge projections is clearly
 780 expressed in the near future, these differences will dis-
 781 appear in the far future, likely due to increasing climate vari-
 782 ability. Similar findings were also reported by Hattermann
 783 et al. (2018), who determined that uncertainty associated
 784 with GCMs is most pronounced during the seasons and
 785 in the regions where the river flow regime is dominated
 786 by precipitation. Many studies accomplished using CMIP6
 787 climate forcing models and scenarios (Wen et al., 2021;
 788 Haider et al., 2023; Núñez Mejía et al., 2023) as well as
 789 those employing CMIP5 tools (Tian et al., 2016; Su et al.,
 790 2017; Vetter et al., 2017; Senatore et al., 2022; Jeantet
 791 et al., 2023) have demonstrated that the choice of repre-
 792 sentative GCMs has a significant impact on the outcomes
 793 of climate impact assessments. In many cases, the domi-
 794 nant source of uncertainty in modeling results stems from
 795 the choice of GCMs rather than from the selection of emis-
 796 sion scenarios. Moreover, different techniques selected
 797 for the ranking procedure may produce different sets of
 798 suitable GCMs for the studied river catchments. Although
 799 there is no universally accepted method for ranking GCMs,
 800 and the process remains inherently subjective (Anil et al.,
 801 2021), it can still be an excellent way to reduce uncertainty
 802 in the final result (Rahman and Pekkak, 2024). The high
 803 GCM-related uncertainty poses significant challenges for
 804 decision-makers and water resource managers, making it
 805 difficult to develop robust adaptation strategies. Projected
 806 changes in runoff patterns affected by this uncertainty may
 807 have serious implications for water availability, ecosystem
 808 health, agriculture, and flood risk management in affected
 809 regions. Therefore, further research is needed to better
 810 understand the sources of GCM uncertainty (Hattermann
 811 et al., 2018) and to improve selection methodologies. This
 812 will ultimately enhance the robustness of climate change

813 impact assessments and support more effective policymak-
 814 ing.

5. Conclusions

816 This study developed and applied a ranking procedure
 817 based on five criteria to identify the best-performing GCMs,
 818 thereby enhancing the reliability of runoff projections.
 819 Based on this approach, three climate models – MIROC-
 820 ES2L, CMCC-CM2-SR5, and KACE-1-0-G – were identified
 821 as best representing Lithuania's climatic conditions. Ac-
 822 cording to the selected GCMs, significant future changes
 823 in air temperature and precipitation were estimated. Tem-
 824 peratures were projected to rise by up to 2.8°C in the near
 825 future and up to 6.4°C in the far future, with the most
 826 pronounced seasonal increases occurring in winter and
 827 summer. Changes in annual precipitation were relatively
 828 modest, with increases up to 5%. Seasonal variability
 829 was anticipated to be greater, with winter precipitation
 830 potentially increasing by as much as 33.7% and summer
 831 precipitation decreasing by up to 23.2%, depending on the
 832 region and scenario. Runoff projections revealed a sub-
 833 stantial decline, with an average annual runoff decreasing
 834 by 12–24% in the near future and 15–42% in the far fu-
 835 ture, relative to the reference period. Notably, low flow
 836 conditions (Q95) were projected to diminish by approxi-
 837 mately two-thirds in the far future, posing critical risks for
 838 hydrological regimes. The uncertainty assessment high-
 839 lighted that selected GCMs contributed up to two-thirds
 840 of the total uncertainty, confirming the utility of the rank-
 841 ing method for model selection in the absence of regional
 842 climate models.

843 Despite limitations due to low climate model resolu-
 844 tion, this study improves our understanding of future low-
 845 land river runoff changes. The use of newly developed
 846 regional climate models will likely enhance the accuracy
 847 of Lithuanian runoff projections.

Acknowledgements

848 The authors wish to thank the Lithuanian Hydrometeo-
 849 rological Service under the Ministry of Environment for
 850 providing the daily hydrometeorological data.

Funding

852 This research received no external funding.

Data availability statement

854 All raw data utilized in this study belong to the respective
 855 institutions listed in the "Study area and data" section. To
 856 gain access to these data, please direct a justified request
 857 to the relevant institutions.

Conflict of interest

859 None declared.

References

861 Afzal, M., Vavlas, N., Ragab, R., 2021. *Modelling study to*
 862 *quantify the impact of future climate and land use*
 863 *changes on water resources availability at catchment*
 864 *scale*. *J. Water Clim. Change* 12, 339–361.
 865 <https://doi.org/10.21166/wcc.2020.117>

866 Akstinas, V., Jakimavičius, D., Meilutytė-Lukauskienė, D.,
 867 Kriauciūnienė, J., Šarauskienė, D., 2020. *Uncertainty*
 868 *of annual runoff projections in Lithuanian rivers under*
 869 *a future climate*. *Hydrol. Res.* 51, 257–271.
 870 <https://doi.org/10.21166/nh.2019.004>

871 Akstinas, V., Šarauskienė, D., Kriauciūnienė, J., Nazarenko,
 872 S., Jakimavičius, D., 2022. *Spatial and Temporal Changes*
 873 *in Hydrological Regionalization of Lowland Rivers*. *Int.*
 874 *J. Environ. Res.* 16, 1.
 875 <https://doi.org/10.1007/s41742-021-00380-8>

876 Alan Yeakley, J., Ervin, D., Chang, H., Granek, E.F., Dujon,
 877 V., Shandas, V., Brown, D., 2016. *Ecosystem services*
 878 *of streams and rivers*. [In:] Gilvear, D.J., Greenwood,
 879 M.T., Thoms, M.C., Wood, P.J. (Eds.), *River Science*, Wiley,
 880 335–352.
 881 <https://doi.org/10.1002/9781118643525.ch17>

882 Anil, S., P. A.R., Vema, V.K., 2024. *Catchment response to*
 883 *climate change under CMIP6 scenarios: a case study*
 884 *of the Krishna River Basin*. *J. Water Clim. Change* 15,
 885 476–498.
 886 <https://doi.org/10.21166/wcc.2024.442>

887 Bergström, S., 1992. *The HBV model – its structure and*
 888 *applications*. SMHI RH No 4. Norrköping, 32.

889 Bhagine, V., Lee, H.S., Cabrera, J.S., Kubota, T., Pradana, R.P.,
 890 Fajary, F.R., Nimiya, H., 2024. *Identification of optimal*
 891 *CMIP6 GCMs for future typical meteorological year in*
 892 *major cities of Indonesia using multi-criteria decision*
 893 *analysis*. *Front. Environ. Sci.* 12, 1341807.
 894 <https://doi.org/10.3389/fenvs.2024.1341807>

895 Bian, G., Zhang, J., Chen, J., Song, M., He, R., Liu, C., Liu, Y.,
 896 Bao, Z., Lin, Q., Wang, G., 2021. *Projecting Hydrological*
 897 *Responses to Climate Change Using CMIP6 Climate*
 898 *Scenarios for the Upper Huai River Basin, China*. *Front.*
 899 *Environ. Sci.* 9, 759547.
 900 <https://doi.org/10.3389/fenvs.2021.759547>

901 Bongaarts, J., 2019. *IPBES, 2019. Summary for policy*
 902 *makers of the global assessment report on biodiversity*
 903 *and ecosystem services of the Intergovernmental*
 904 *Science-Policy Platform on Biodiversity and Ecosystem*
 905 *Services*. *Popul. Dev. Rev.* 45, 6800–681.
 906 <https://doi.org/10.1111/padr.12283>

907 Brêda, J.P.L.F., De Paiva, R.C.D., Collischon, W., Bravo, J.M.,
 908 Siqueira, V.A., Steinke, E.B., 2020. *Climate change im-*
 909 *pacts on South American water balance from a continental-*
 910 *scale hydrological model driven by CMIP5 projections*.
 911 *Clim. Change* 159, 503–522.
 912 <https://doi.org/10.1007/s10584-020-02667-9>

913 Calvin, K., Dasgupta, D., Krinner, G., Mukherji, A., Thorne, P.,
 914 W., Trisos, C., Romero, J., et al., 2023. *Climate Change*
 915 *2023: Synthesis Report. Contribution of Working Groups*
 916 *I, II and III to the Sixth Assessment Report of the In-*
 917 *tergovernmental Panel on Climate Change* First Inter-
 918 *governmental Panel on Climate Change*, IPCC, Geneva,
 919 *Switzerland*.
 920 <https://doi.org/10.59327/IPCC/AR6-9789291691647>

921 Chen, H., Sun, J., Lin, W., Xu, H., 2020. *Comparison of CMIP6*
 922 *and CMIP5 models in simulating climate extremes*. *Sci-*
 923 *ence Bull.* 65, 1415–1418.
 924 <https://doi.org/10.1016/j.scib.2020.05.015>

925 Clark, M.P., Bierkens, M.F.P., Samaniego, L., Woods, R.A., Uij-
 926 jenhoet, R., Bennet, K.E., Pauwels, V.R.N., Cai, X., Wood,
 927 A.W., Peters-Lidard, C.D., 2017. *The evolution of process-*
 928 *based hydrologic models: Historical challenges and the*
 929 *collective quest for physical realism*.
 930 <https://doi.org/10.5194/hess-2016-693>

931 Copernicus Climate Change Service (C3S), 2024. *European*
 932 *State of the Climate 2023. Copernicus Climate Change*
 933 *Service (C3S)*.
 934 <https://doi.org/10.24381/BS9V-8C66>

935 Dallison, R.J.H., Williams, A.P., Harris, I.M., Patil, S.D., 2022.
 936 *Modelling the impact of future climate change on stream-*
 937 *flow and water quality in Wales, UK*. *Hydrol. Sci. J.* 67,
 938 939–962.
 939 <https://doi.org/10.1080/02626667.2022.2044045>

940 Dialogue Earth, 2022. *Irreversible climate impacts already*
 941 *here: IPCC*. <https://dialogue.earth/en/climate/ipcc-sixth-assessment-stark-warning/> (accessed on
 942 04-06-2024)

943 Döll, P., Jiménez-Cisneros, B., Oki, T., Arnell, N.W., Benito, G.,
 944 Cogley, J.G., Jiang, T., Kundzewicz, Z.W., Mwakalila, S.,
 945 Nishijima, A., 2015. *Integrating risks of climate change*
 946 *into water management*. *Hydrol. Sci. J.* 60, 4–13.
 947 <https://doi.org/10.1080/02626667.2014.967250>

948 Donnelly, C., Greuell, W., Andersson, J., Gerten, D., Pisacane,
 949 G., Roudier, P., Ludwig, F., 2017. *Impacts of climate*
 950 *change on European hydrology at 1.5, 2 and 3 degrees*
 951 *mean global warming above preindustrial level*. *Clim.*
 952 *Change* 143, 13–26.
 953 <https://doi.org/10.1007/s10584-017-1971-7>

954 Duan, K., Sun, G., McNulty, S.G., Caldwell, P.V., Cohen, E.C.,
 955 Sun, S., Aldridge, H.D., Zhou, D., Zhang, L., Zhang, Y.,
 956 2017. *Future shift of the relative roles of precipitation*
 957 *and temperature in controlling annual runoff in the*
 958 *conterminous United States*. *Hydrol. Earth Syst. Sci.*
 959 21, 5517–5529.
 960 <https://doi.org/10.5194/hess-21-5517-2017>

961 Etukudoh, E.A., Ilojianya, V. I., Ayorinde, O., B., Daudu, C.,
 962 D., Adefemi, A., Hamdan, A., 2024. *Review of climate*
 963 *change impact on water availability in the USA and*
 964 *Africa*. *Int. J. Sci. Res. Arch.* 11, 942–951.
 965 <https://doi.org/10.30574/ijrsa.2024.11.1.0169>

966 Eyring, V., Bony, S., Meehl, G.A., Senior, C.A., Stevens, B.,
 967 Stouffer, R.J., Taylor, K.E., 2016. *Overview of the Cou-*
 968 *pled Model Intercomparison Project Phase 6 (CMIP6)*
 969 <https://doi.org/10.5194/hess-2016-693>

971 *experimental design and organization.* Geosci. Model
 972 Dev. 9, 1937–1958.
 973 <https://doi.org/10.5194/gmd-9-1937-2016> 1026

974 European Environment Agency (EEA), 2023. *Economic
 975 losses from weather- and climate-related extremes in
 976 Europe.* <https://www.eea.europa.eu/en/analysis/indicators/economic-losses-from-climate-related> 1027

977 (accessed on 02-06-2024) 1028

978 Gailiušis, B., Jablonskis, J., Kovalenkoviene, M., 2001. *Lithuanian
 979 rivers: hydrography and runoff.* Lithuanian Energy Inst., Kaunas, Lithuania. 792 pp. 1029

980 Gebrechorkos, S., Leyland, J., Slater, L., Wortmann, M., Ash-
 981 worth, P.J., Bennett, G.L., Boothroyd, R., Cloke, H., De-
 982 lorme, P., Griffith, H., Hardy, R., Hawker, L., McLelland,
 983 S., Neal, J., Nicholas, A., Tatem, A.J., Vahidi, E., Parsons,
 984 D.R., Darby, S.E., 2023. *A high-resolution daily global
 985 dataset of statistically downscaled CMIP6 models for
 986 climate impact analyses.* Sci. Data 10, 611. 1030

987 <https://doi.org/10.1038/s41597-023-02528-x> 1031

988 Gebresellase, S.H., Wu, Z., Xu, H., Muhammad, W.I., 2022. 1032

989 *Evaluation and selection of CMIP6 climate models in
 990 Upper Awash Basin (UBA), Ethiopia: Evaluation and
 991 selection of CMIP6 climate models in Upper Awash Basin
 992 (UBA), Ethiopia.* Theor. Appl. Climatol. 149, 1521 1033

993 –1547.
 994 <https://doi.org/10.1007/s00704-022-04056-x> 1034

995 Grose, M.R., Narsey, S., Delage, F.P., Dowdy, A.J., Bador, M.,
 996 Boschat, G., Chung, C., Kajtar, J.B., Rauniyar, S., Fre-
 997 und, M.B., Lyu, K., Rashid, H., Zhang, X., Wales, S., Tren-
 998 ham, C., Holbrook, N.J., Cowan, T., Alexander, L., Ar-
 999 blaster, J.M., Power, S., 2020. *Insights From CMIP6
 1000 for Australia's Future Climate.* Earth's Futur. 8, 1050

1001 e2019EF001469.
 1002 <https://doi.org/10.1029/2019EF001469> 1051

1003 Guan, X., Zhang, J., Bao, Z., Liu, C., Jin, J., Wang, G., 2021. 1052

1004 *Past variations and future projection of runoff in typical
 1005 basins in 10 water zones, China.* Sci. Total Environ. 798,
 1006 149277.
 1007 <https://doi.org/10.1016/j.scitotenv.2021.149277> 1053

1008 Gusain, A., Ghosh, S., Karmakar, S., 2020. *Added value of
 1009 CMIP6 over CMIP5 models in simulating Indian summer
 1010 monsoon rainfall.* Atmos. Res. 232, 104680. 1054

1011 <https://doi.org/10.1016/j.atmosres.2019.104680> 1055

1012 Haider, S., Masood, M.U., Rashid, M., Alshehri, F., Pande, C.B.,
 1013 Katipoğlu, O.M., Costache, R., 2023. *Simulation of the
 1014 Potential Impacts of Projected Climate and Land Use
 1015 Change on Runoff under CMIP6 Scenarios.* Water 15,
 1016 3421.
 1017 <https://doi.org/10.3390/w15193421> 1056

1018 Hattermann, F.F., Vetter, T., Breuer, L., Su, B., Daggupati, P.,
 1019 Donnelly, C., Fekete, B., Flörke, F., Gosling, S.N., Hoff-
 1020 mann, P., Liersch, S., Masaki, Y., Motovilov, Y., Müller, C.,
 1021 Samaniego, L., Stacke, T., Wada, Y., Yang, T., Krysnaova,
 1022 V., 2018. *Sources of uncertainty in hydrological climate
 1023 impact assessment: a cross-scale study.* Environ. Res.
 1024 Lett. 13, 015006.
 1025 <https://doi.org/10.1088/1748-9326/aa9938> 1027

1026 Held, I.M., Soden, B.J., 2006. *Robust Responses of the Hydro-
 1027 logical Cycle to Global Warming.* J. Clim. 19, 5686–5699.
 1028 <https://doi.org/10.1175/JCLI3990.1> 1029

1029 Integrated Hydrological Modelling System (IHMS), 2005. 1030

1030 *Manual. Version 5.8.* Swedish Meteorol. Hydrol.Inst.,
 1031 115 pp.
 1032 Intergovernmental Panel On Climate Change (IPCC), 2023a. 1033

1033 *Climate Change 2022 – Impacts, Adaptation and Vul-
 1034 nerability: Working Group II Contribution to the Sixth
 1035 Assessment Report of the Intergovernmental Panel on
 1036 Climate Change.* 1st edn., Cambridge Univ. Press.
 1037 <https://doi.org/10.1017/9781009325844> 1038

1038 Intergovernmental Panel On Climate Change (IPCC), 2023b. 1039

1039 *Climate Change 2021 – The Physical Science Basis: Work-
 1040 ing Group I Contribution to the Sixth Assessment Report
 1041 of the Intergovernmental Panel on Climate Change.* 1st
 1042 edn., Cambridge Univ. Press.
 1043 <https://doi.org/10.1017/9781009157896> 1044

1044 Iqbal, Z., Shahid, S., Ahmed, K., Ismail, T., Ziarh, G.F., Chung,
 1045 E.-S., Wang, X., 2021. *Evaluation of CMIP6 GCM rainfall
 1046 in mainland Southeast Asia.* Atmos. Res. 254, 105525.
 1047 <https://doi.org/10.1016/j.atmosres.2021.105525> 1048

1048 Jakimavičius, D., Adžgauskas, G., Šarauskienė, D.,
 1049 Kriauciūnienė, J., 2020. *Climate Change Impact on Hy-
 1050 dropower Resources in Gauged and Ungauged Lithua-
 1051 nian River Catchments.* Water 12, 3265.
 1052 <https://doi.org/10.3390/w12113265> 1053

1053 Jakimavičius, D., Kriauciūnienė, J., Šarauskienė, D., 2018. 1054

1054 *Impact of climate change on the Curonian Lagoon water
 1055 balance components, salinity and water temperature
 1056 in the 21st century.* Oceanologia 60 (3), 378–389.
 1057 <https://doi.org/10.1016/j.oceano.2018.02.003> 1058

1058 Jeantet, A., Thirel, G., Lemaitre-Basset, T., Tournebize, J.,
 1059 2023. *Uncertainty propagation in a modelling chain
 1060 of climate change impact for a representative French
 1061 drainage site.* Hydrol. Sci. J. 68, 1426–1442.
 1062 <https://doi.org/10.1080/02626667.2023.2203322> 1063

1063 Jiang, T., Su, B., Huang, J., Zhai, J., Xia, J., Tao, H., Wang, Y.,
 1064 Sun, H., Luo, Y., Zhang, L., Wang, G., Zhan, C., Xiong,
 1065 M., Kundzewicz, Z.W., 2020. *Each 0.5°C of Warming
 1066 Increases Annual Flood Losses in China by More than
 1067 US\$60 Billion.* Bull. Am. Meteorol. Soc. 101, E1464
 1068 –E1474.
 1069 <https://doi.org/10.1175/BAMS-D-19-0182.1> 1070

1070 Kim, Y.-H., Min, S.-K., Zhang, X., Sillmann, J., Sandstad, M.,
 1071 2020. *Evaluation of the CMIP6 multi-model ensemble
 1072 for climate extreme indices.* Weather Clim. Extremes
 1073 29, 100269.
 1074 <https://doi.org/10.1016/j.wace.2020.100269> 1075

1075 Kis, A., Pongrácz, R., 2024. *The projected changes of hydro-
 1076 logical indicators in European catchments with differ-
 1077 ent climatic conditions.* Hydrol. Sci. J. 69, 1797–1812.
 1078 <https://doi.org/10.1080/02626667.2024.2390908> 1079

1079

1081 Kriauciuniene, J., Jakimavicius, D., Sarauskiene, D., Kaliatka,
 1082 T., 2013. *Estimation of uncertainty sources in the pro-*
 1083 *jections of Lithuanian river runoff.* Stoch. Environ. Res.
 1084 Risk. Assess. 27, 769–784.
 1085 <https://doi.org/10.1007/s00477-012-0608-7> 1136

1086 Kriauciūnienė, J., Meilutytė-Barauskienė, D., Rimkus, E.,
 1087 Kažys, J., Vincevičius, A., 2008. *Climate change impact*
 1088 *on hydrological processes in Lithuanian Nemunas river*
 1089 *basin.* Baltica 21, 51–61.

1090 Kriauciūnienė, J., Virbickas, T., Šarauskienė, D., Jakimav-
 1091 ičius, D., Kažys, J., Bukanis, A., Kesminas, V., Povilaitis,
 1092 A., Dainys, J., Akstinas, V., Jurgelėnaitė, A., Meilutytė-
 1093 Lukauskienė, D., Tomkevičienė, A., 2019. *Fish assem-
 1094 blages under climate change in Lithuanian rivers.* Sci.
 1095 Total Environ. 661, 563–574.
 1096 <https://doi.org/10.1016/j.scitotenv.2019.01.142> 1141

1097 Kundzewicz, Z.W., Krysanova, V., Benestad, R.E., Hov, Ø.,
 1098 Piniewski, M., Otto, I.M., 2018. *Uncertainty in climate*
 1099 *change impacts on water resources.* Environ. Sci. Policy
 1100 79, 1–8.
 1101 <https://doi.org/10.1016/j.envsci.2017.10.008> 1142

1102 Kurniadi, A., Weller, E., Kim, Y., Min, S., 2023. *Evaluation of*
 1103 *Coupled Model Intercomparison Project Phase 6 model-sim-
 1104 ulated extreme precipitation over Indonesia.* Int. J. Cli-
 1105 matol. 43, 174–196.
 1106 <https://doi.org/10.1002/joc.7744> 1143

1107 Lane, R.A., Kay, A.L., 2021. *Climate Change Impact on the*
 1108 *Magnitude and Timing of Hydrological Extremes Across*
 1109 *Great Britain.* Front. Water 3, 684982.
 1110 <https://doi.org/10.3389/frwa.2021.684982> 1144

1111 Lennox, R.J., Crook, D.A., Moyle, P.B., Struthers, D.P., Cooke,
 1112 S.J., 2019. *Toward a better understanding of freshwater*
 1113 *fish responses to an increasingly drought-stricken world.*
 1114 Rev. Fish. Biol. Fisheries 29, 71–92.
 1115 <https://doi.org/10.1007/s11160-018-09545-9> 1145

1116 Martel, J.-L., Brissette, F., Troin, M., Arsenault, R., Chen, J.,
 1117 Su, T., Lucas-Picher, P., 2022. *CMIP5 and CMIP6 Model*
 1118 *Projection Comparison for Hydrological Impacts Over*
 1119 *North America.* Geophys. Res. Lett. 49, e2022GL098364.
 1120 <https://doi.org/10.1029/2022GL098364> 1146

1121 Meinshausen, M., Nicholls, Z.R.J., Lewis, J., Gidden, M.J., Vo-
 1122 gel, E., Freund, M., Beyerle, U., Gessner, C., Nauels, A.,
 1123 Bauer, N., Canadell, J.G., Daniel, J.S., John, A., Krum-
 1124 mel, P.B., Luderer, G., Meinshausen, N., Montzka, S.A.,
 1125 Rayner, P.J., Reimann, S., Smith, S.J., Van Den Berg, M.,
 1126 Velders, G.J.M., Vollmer, M.K., Wang, R.H.J., 2020. *The*
 1127 *shared socio-economic pathway (SSP) greenhouse gas*
 1128 *concentrations and their extensions to 2500.* Geosci.
 1129 Model Dev. 13, 3571–3605.
 1130 <https://doi.org/10.5194/gmd-13-3571-2020> 1147

1131 Muelchi, R., Rössler, O., Schwanbeck, J., Weingartner, R.,
 1132 Martius, O., 2021. *River runoff in Switzerland in a chang-
 1133 ing climate – runoff regime changes and their time of*
 1134 *emergence.* Hydrol. Earth Syst. Sci. 25, 3071–3086.
 1135 <https://doi.org/10.5194/hess-25-3071-2021> 1148

1136 Murphy, C., Kettle, A., Meresa, H., Golian, S., Bruen, M.,
 1137 O'Loughlin, F., Mellander, P.-E., 2023. *Climate Change*
 1138 *Impacts on Irish River Flows: High Resolution Scenarios*
 1139 *and Comparison with CORDEX and CMIP6 Ensembles.*
 1140 Water Resour. Manage. 37, 1841–1858.
 1141 <https://doi.org/10.1007/s11269-023-03458-4> 1142

1142 Nazarenko, S., Meilutytė-Lukauskienė, D., Šarauskienė, D.,
 1143 Kriauciūnienė, J., 2022. *Spatial and Temporal Patterns*
 1144 *of Low-Flow Changes in Lowland Rivers.* Water 14, 801.
 1145 <https://doi.org/10.3390/w14050801> 1146

1146 Nazarenko, S., Šarauskienė, D., Putrenko, V., Kriauciūnienė,
 1147 J., 2023. *Evaluating Hydrological Drought Risk in Lithua-
 1148 nia.* Water 15, 2830.
 1149 <https://doi.org/10.3390/w15152830> 1150

1150 Nguyen-Duy, T., Ngo-Duc, T., Desmet, Q., 2023. *Performance*
 1151 *evaluation and ranking of CMIP6 global climate models*
 1152 *over Vietnam.* J. Water Clim. Change 14, 1831–1846.
 1153 <https://doi.org/10.2166/wcc.2023.454> 1154

1154 Núñez Mejía, S.X., Mendoza Paz, S., Tabari, H., Willems, P.,
 1155 2023. *Climate change impacts on hydrometeorologi-
 1156 cal and river hydrological extremes in Quito, Ecuador.*
 1157 J. Hydrol.: Reg. Stud. 49, 101522.
 1158 <https://doi.org/10.1016/j.ejrh.2023.101522> 1159

1159 Palmer, T.E., Booth, B.B.B., McSweeney, C.F., 2021. *How does*
 1160 *the CMIP6 ensemble change the picture for European*
 1161 *climate projections?* Environ. Res. Lett. 16, 094042.
 1162 <https://doi.org/10.1088/1748-9326/ac1ed9> 1163

1163 Pervin, L., Gan, T.Y., Scheepers, H., Islam, M.S., 2021. *Appli-
 1164 cation of the HBV model for the future projections of wa-
 1165 ter levels using dynamically downscaled global climate*
 1166 *model data.* J. Water Clim. Change 12, 2364–2377.
 1167 <https://doi.org/10.2166/wcc.2021.302> 1168

1168 Piniewski, M., Szcześniak, M., Huang, S., Kundzewicz, Z.W.,
 1169 2018. *Projections of runoff in the Vistula and the Odra*
 1170 *river basins with the help of the SWAT model.* Hydrol.
 1171 Res. 49, 303–317.
 1172 <https://doi.org/10.2166/nh.2017.280> 1173

1173 Rahman, A., Pekkat, S., 2024. *Identifying and ranking of*
 1174 *CMIP6-global climate models for projected changes in*
 1175 *temperature over Indian subcontinent.* Sci. Rep. 14,
 1176 3076.
 1177 <https://doi.org/10.1038/s41598-024-52275-1> 1178

1178 Raju, K.S., Kumar, D.N., 2020. *Review of approaches for se-
 1179 lection and ensembling of GCMs.* J. Water Clim. Change
 1180 11, 577–599.
 1181 <https://doi.org/10.2166/wcc.2020.128> 1182

1182 Ritter, A., Muñoz-Carpena, R., 2013. *Performance evalua-
 1183 tion of hydrological models: Statistical significance*
 1184 *for reducing subjectivity in goodness-of-fit assessments.*
 1185 J. Hydrol. 480, 33–45.
 1186 <https://doi.org/10.1016/j.jhydrol.2012.12.004> 1187

1187 Rivera, P., 2024. *Climate change projections in Guatemala:*
 1188 *temperature and precipitation changes according to*
 1189 *CMIP6 models.* Model. Earth Syst. Environ. 10, 2031
 1190

1191 -2049.
 1192 <https://doi.org/10.1007/s40808-023-01881-5>

1193 Šarauskienė, D., Akstinas, V., Kriauciūnienė, J., Jakimavičius,
 1194 D., Bukantis, A., Kažys, J., Povilaitis, A., Ložys, L., Kesmi-
 1195 nės, V., Virbickas, T., Pliuraitė, V., 2018. *Projection of*
 1196 *Lithuanian river runoff, temperature and their extremes*
 1197 *under climate change*. *Hydrol. Res.* 49, 344–362.
 1198 <https://doi.org/10.2166/nh.2017.007>

1199 Šarauskienė, D., Akstinas, V., Nazarenko, S., Kriauciūnienė,
 1200 J., Jurgelėnaitė, A., 2020. *Impact of physico-geographical*
 1201 *factors and climate variability on flow intermittency*
 1202 *in the rivers of water surplus zone*. *Hydrol. Proc.* 34,
 1203 4727–4739.
 1204 <https://doi.org/10.1002/hyp.13912>

1205 Senatore, A., Fuoco, D., Maiolo, M., Mendicino, G., Smi-
 1206 atek, G., Kunstmann, H., 2022. *Evaluating the uncer-*
 1207 *tainty of climate model structure and bias correction*
 1208 *on the hydrological impact of projected climate change*
 1209 *in a Mediterranean catchment*. *J. Hydrol.: Reg. Stud.*
 1210 42, 101120.
 1211 <https://doi.org/10.1016/j.ejrh.2022.101120>

1212 Sleziaik, P., Výleta, R., Hlavčová, K., Danáčová, M., Aleksić, M.,
 1213 Szolgyay, J., Kohnová, S., 2021. *A Hydrological Modeling*
 1214 *Approach for Assessing the Impacts of Climate Change*
 1215 *on Runoff Regimes in Slovakia*. *Water* 13, 3358.
 1216 <https://doi.org/10.3390/w13233358>

1217 Stonevičius, E., Rimkus, E., Štaras, A., Kažys, J., Valiuške-
 1218 vičius, G., 2017. *Climate change impact on the Nemunas*
 1219 *River basin hydrology in the 21st century*. *Boreal*
 1220 *Environ. Res.* 22, 49–65.

1221 Su, B., Huang, J., Zeng, X., Gao, C., Jiang, T., 2017. *Impacts*
 1222 *of climate change on streamflow in the upper Yangtze*
 1223 *River basin*. *Clim. Change* 141, 533–546.
 1224 <https://doi.org/10.1007/s10584-016-1852-5>

1225 Susnik, J., Masia, S., Teutschbein, C., 2023. *Water as a key*
 1226 *enabler of nexus systems (water-energy-food)*. *Camb.*
 1227 *Prisms Water* 1–29.
 1228 <https://doi.org/10.1017/wat.2023.1>

1229 Tariq, M., Rohith, A.N., Cibin, R., Aruffo, E., Abouzied, G.A.A.,
 1230 Carlo, P.D., 2024. *Understanding future hydrologic chal-*
 1231 *lenges: Modelling the impact of climate change on river*
 1232 *runoff in central Italy*. *Environ. Challenges* 15, 100899.
 1233 <https://doi.org/10.1016/j.envc.2024.100899>

1234 The Guardian, 2021. *Major climate changes inevitable and*
 1235 *irreversible – IPCC's starker warning yet*. [https://ww-
 1239 heguardian.com/science/2021/aug/09/huma-
 1240 ns-have-caused-unprecedented-and-irreversibl-
 1241 e-change-to-climate-scientists-warn](https://ww-

 1236 w.theguardian.com/science/2021/aug/09/huma-

 1237 ns-have-caused-unprecedented-and-irreversibl-

 1238 e-change-to-climate-scientists-warn) (accessed on
 1242 28-05-2024)

1243 Tian, Y., Xu, Y.-P., Booij, M.J., Cao, L., 2016. *Impact assess-*
 1244 *ment of multiple uncertainty sources on high flows un-*
 1245 *der climate change*. *Hydrol. Res.* 47, 61–74.
 1246 <https://doi.org/10.2166/nh.2015.008>

1247 Vetter, T., Reinhardt, J., Flörke, M., Van Griensven, A., Hatter-
 1248 mann, F., Huang, S., Koch, H., Pechlivanidis, I.G., Plötner,
 1249 S., Seidou, O., Su, B., Vervoort, R.W., Krysanova, V., 2017. *Evaluation of sources of uncertainty in projected hydro-*
 1250 *logical changes under climate change in 12 large-scale*
 1251 *river basins*. *Clim. Change* 141, 419–433.
 1252 <https://doi.org/10.1007/s10584-016-1794-y>

1253 Wei, L., Xin, X., Li, Q., Wu, Y., Tang, H., Li, Y., Yang, B., 2023. *Simulation and projection of climate extremes in China*
 1254 *by multiple Coupled Model Intercomparison Project*
 1255 *Phase 6 models*. *Int. J. Climatol.* 43, 219–239.
 1256 <https://doi.org/10.1002/joc.7751>

1257 Wen, K., Gao, B., Li, M., 2021. *Quantifying the Impact of Fu-*
 1258 *ture Climate Change on Runoff in the Amur River Basin*
 1259 *Using a Distributed Hydrological Model and CMIP6 GCM*
 1260 *Projections*. *Atmosphere* 12, 1560.
 1261 <https://doi.org/10.3390/atmos12121560>

1262 Woodward, G., Perkins, D.M., Brown, L.E., 2010. *Climate*
 1263 *change and freshwater ecosystems: impacts across mul-*
 1264 *tiple levels of organization*. *Phil. Trans. R. Soc. B* 365,
 1265 2093–2106.
 1266 <https://doi.org/10.1098/rstb.2010.0055>

1267 Xin, X., Wu, T., Zhang, J., Yao, J., Fang, Y., 2020. *Comparison of*
 1268 *CMIP6 and CMIP5 simulations of precipitation in China*
 1269 *and the East Asian summer monsoon*. *Int. J. Climatol.*
 1270 40, 6423–6440.
 1271 <https://doi.org/10.1002/joc.6590>

1272 Xiong, J., Guo, S., Abhishek, Chen, J., Yin, J., 2022. *Global*
 1273 *evaluation of the “dry gets drier, and wet gets wetter”*
 1274 *paradigm from a terrestrial water storage change per-*
 1275 *spective*. *Hydrol. Earth Syst. Sci.* 26, 6457–6476.
 1276 <https://doi.org/10.5194/hess-26-6457-2022>

1277 Yang, H., Zhou, F., Piao, S., Huang, M., Chen, A., Ciais, P.,
 1278 Li, Y., Lian, X., Peng, S., Zeng, Z., 2017. *Regional pat-*
 1279 *terns of future runoff changes from Earth system models*
 1280 *constrained by observation*. *Geophys. Res. Lett.* 44,
 1281 5540–5549.
 1282 <https://doi.org/10.1002/2017GL073454>

1283 Yang, T., Ding, J., Liu, D., Wang, X., Wang, T., 2019. *Combined*
 1284 *Use of Multiple Drought Indices for Global Assessment of*
 1285 *Dry Gets Drier and Wet Gets Wetter Paradigm*. *J. Climate*
 1286 32, 737–748.
 1287 <https://doi.org/10.1175/JCLI-D-18-0261.1>